Estimating Crustal Velocity Structure in Alaska from Acoustic-to-Seismic Coupling from the 2022 Hunga Eruption, Tonga

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Cleared for release

Motivation

Use the once-in-a-lifetime source from the Hunga, Tonga eruption to examine crustal structure using acoustic-seismic coupling at deeper depths than is typically possible

Tonga Geological Survey

Hunga, Tonga Infrasound in Alaska

- 150 stations equipped with colocated, broadband:
	- seismic (BH?)
	- infrasound (BDF)
	- barometer (BDO)
- **Large pressure amplitudes** (> 60.0 Pa) at huge offsets!

— Sorrels, 1971

Pressure-to-Seismic Coupling

Infrasound (BDF) Seismic (BHZ)

10250

10000

9750

9500 **a** tanc

9250

9000

8750

8500

Έ

— Sorrels, 1971

Pressure-to-Seismic Coupling

Infrasound (BDF) Seismic (BHZ)

14:30

10250

10000

9750

9500

9250

9000

8750

8500

15:00

— Sorrels, 1971

Pressure-to-Seismic Coupling

— Sorrels, 1971

Pressure-to-Seismic Coupling

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Pressure-to-Seismic Coupling

Seismic (BHZ) Infrasound (BDF)

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a tanc

"*...colocated data allow us to identify the frequency range of strong coupling between the atmosphere and the solid Earth and provide us information on how the solid Earth responds to surface pressure changes.*" —Tanimoto and Wang, 2018

Pressure-to-Seismic Coupling

Seismic (BHZ) Infrasound (BDF)

We choose bands with good coherence (> 0.8) between the seismic and pressure:

The magnitude squared coherence between the pressure and vertical seismic is:

$$
\gamma^2 = \frac{\overline{G_{PS}}^2}{\overline{G_{PP}}\,\overline{G_{SS}}}
$$

Where G_{pS} is the cross spectral density of the pressure and seismic, and G_{pp} and G_{SS} are the autospectral densities

Network Coherence

Broadband infrasound and seismic

Network Coherence

Broadband infrasound and seismic

Network Coherence

Broadband infrasound and seismic

Coupling Calculation

Coupling spectra is simply the ratio of seismic amplitudes to infrasound amplitudes:

$$
\Gamma(f) = \sqrt{\frac{PSD_s(f)}{PSD_p(f)}},
$$

Where *PSD s* and *PSD p* are the seismic and pressure power spectral densities, respectively.

Also, a rule of thumb for the depth sensitivity of coupling to material parameters is give by;

 $h = 0.15 \cdot c \cdot T$

Where *h* is depth, *T* is period, and *c* is the pressure source speed. (Tanimoto and Wang, 2019)

coherence (γ^2)

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$$
\Gamma(f) = \sqrt{\frac{PSD_s(f)}{PSD_p(f)}} = \frac{c(\lambda + 2\mu)}{2\mu(\lambda + \mu)}
$$

Where c is sound speed and λ and μ are the first Lame' parameter and rigidity, respectively (Sorrels, 1971)

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Estimating Near-Surface Rigidity from Low-Frequency Noise Using Collocated Pressure and Horizontal Seismic Data Jiong Wang^{*1} and Toshiro Tanimoto

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$$
\bar{\mu} = \frac{\lambda + \mu}{\lambda + 2\mu}\mu
$$

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where $\varrho_{{\mathit{mod}}}$ is a depth-averaged density estimate from CRUST1.0

Compare to existing models:

1.4 Hz (35 m) Model: USGS proxy V_{530} (Allen and Wald, 2007)

0.025 Hz (2.0 km) and 0.00975 Hz (5.0 km) Model: tomographic (**Berg,** *et al.***, 2020)** Depth-averaged to appropriate depth

Results: Mean V_s **for Upper 35 m (~1.4 Hz)**

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Background model from USGS global proxy V_{530}

Results: Mean V_s **for Upper 2.0 km (~0.025 Hz)**

Background tomographic model from Berg, *et al.*, 2020 (depth averages)

Results: Mean V_s **for Upper 5.0 km (~0.00975 Hz)**

Background tomographic model from Berg, *et al.*, 2020 (depth averages)

coherence

Results: Comparison with 2013 Chelyabinsk Bolide

Photo: Alex Alishevskikh

 -120 $\frac{100}{5}$
 $\frac{100}{5}$

 -140 -160

 $-180E$

Conclusions:

- **● Pressure waves from the Hunga,Tonga eruption produced air-to-ground coupled waves that were beautifully recorded in Alaska**
- **● Microseismic amplitudes generally exceed coupled seismic amplitudes**
- **● Coupling was inversely proportional to the rigidity of the elastic medium, and this was used to estimate bulk** V_s
- **● Successful estimate of mean** *V s* **to depths of:**
	- **○ 35 m**
	- **○ 2.0 km**
	- **○ 5.0 km**

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Thank You! Questions or Comments?

