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Introduction: detection capability and transmission losses

detection capability with infrasound technology builds upon relevant estimations of

transmission losses at the surface across the network throughout the year (Figure 1).

January 1 April 1 April 1 April 1 April 1 October 1 October

Figure 1. Smallest signal attenuation expected at 0.8Hz with a 2-station coverage (Le Pichon et al. 2012) using IFS/ECMWF (-60 dB is a factor of 1000 in amplitude.)

- Gravity waves (GW) significantly alter the propagation path of infrasound waves in the middle atmospheric waveguide through partial reflections in the shadow zones (Figure 2), or through the temporary setting of a new stratospheric geometric waveguide.
- However GW are often poorly resolved in atmospheric specifications despite their estimated significant effect on detection thresholds (up to factor of 5-10 on the amplitude) Le Pichon et al. 2019



Figure 2. Explaining infrasound detection of Le Teil's earthquake at Observatoire de Haute Provence (OHP) (Vallage et al. 2021)



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In the literature, there are different ways of investigating/accounting for the impact of GW in infrasound propagation simulations:

- Parameterizations based on the GW universal spectrum e.g. Gardner et al. 1993, as in Vallage et al. 2021
- Stochastic parameterizations accounting for GW intermittency e.g. de la Camara et al., 2015 as in Cugnet et al. 2019
- GW ray-tracing equations applied to a frequency spectrum e.g. Drob et al. 2013
- 3D GW-spectrum model Chunchuzov & Kulichkov, 2019
- Working with high-resolution O(1 km) models explicitly resolving a large part of the GW spectrum
 without using GW parameterizations is another way, given increased computing means.
- We use a dataset of a high-resolution model runs' outputs to demonstrate a method for quantifying the systematic impact of GW across IMS stations, based on transmission losses (TLoss) calculations.

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Data and method: deriving GW perturbations in the stratosphere and building atmospheric specifications



Data	Method	
DYAMOND dataset Stevens et al. 2019, phase II Stephan et al. 2022	GW extraction for field X=U,V,T	
Model : ICON Zängl et al. 2015	 ICON outputs interpolated on a vertical grid 	
Period: 20 Jan. – 29 Feb. 2020 (3-hourly outputs)	with dz=1.5km to match average dz in the	
Initialization : ECMWF/IFS ; freely-running	stratosphere	\sim
Model top : 75km (45 km: avoiding sponge layer)	 background X_{back} obtained by filtering out 	
Configurations:	λ_z < 15 km (3 rd order Butt. filt.) in T	
 dpp0029: dx = 5km (regridded: 0.35° x 0.35°) 	e.g. Baumgarten et al., 2017	OBJECTIVES
 nwp2.5winter: dx = 2.5km (regridded: 0.35° x 0.35°) 	• deriving GW perturbation: $X - X_{back} = \Delta T$	
	Atmospheric specifications for IS	METHODS/DA
Rayleigh lidar observations at OHP e.g. Hauchecorne et al. 1980	Altitude: sticking to 0-45 km only	RESULTS
Observatoire de Haute Provence, France (LATMOS)	\rightarrow avoiding artefacts from upper	CONCLUSIO
Altitude range: 30-90km; Vertical resolution : 75 m;	interpolation with other model	
Accuracy: < 1K (below 70km altitude)	Filtering: only applied in the stratosphere	
Data: 16 night profiles (4 hourly-average)	\rightarrow avoiding filtering out low level jets	
Satellite observations: GRACILE dataset Ern at al 2019	$\begin{bmatrix} At OHP \\ \underbrace{\&} \\ 40 \end{bmatrix} = \begin{bmatrix} 40 \\ 40 \end{bmatrix} = \begin{bmatrix} 40 \\ 40 \end{bmatrix}$	
IR limb sounders HIRDIS (2005-2008) and SAREP (2002-2015)	29 Jan. 2020 gg 20 - 18H gg 20 - 18H gg 20 - 20 -	
\rightarrow zonal averages of En (mean max min) Diat $-2.5^{\circ}/5^{\circ}$		Places de
~ 20 rai averages of EP (mean, max, min), $Diat_{hirdls/saber}=2.5/5$	U (m/s) V (m/s) T (K)	not use thi

Infrasound propagation simulations

Range-independant PE simulations done with NCPAprop *Waxler & Assink, 2019* at OHP and IMS stations **Deriving TLoss differences** between PE simulation using specifications w and w/o GW, respectively

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Results: stratospheric GW across the IMS and impact on transmission loss

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Conclusion: impact of gravity waves (GW) on transmission losses (TLoss)

- We demonstrate a method for quantifying the systematic impact of stratospheric GW across IMS stations, based on TLoss calculations with PE simulations.
- We use a database of state of the art high-resolution model outputs where GW are not parameterized.
- We validate the modelled GW perturbations using Rayleigh lidar data at Observatoire de Haute-Provence.
- We validate the modelled GW amplitudes using satellite products across the IMS based on the GW potential energy (E_p).
- The average impact of GW is much larger at 1 Hz (Tloss increase of up to 40 dB) than at 0.1 Hz (less than 10 dB).
- The impact of GW versus distance-to-station depends on the considered IMS station (hemisphere) with a more or less pronounced impact on the shadow zone. There is no systematic link between GW impact on TLoss and GW energy (latitude). This points at the complex intrication of small-scale structure's role with that of the larger-scale variability (main stratospheric guide).

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